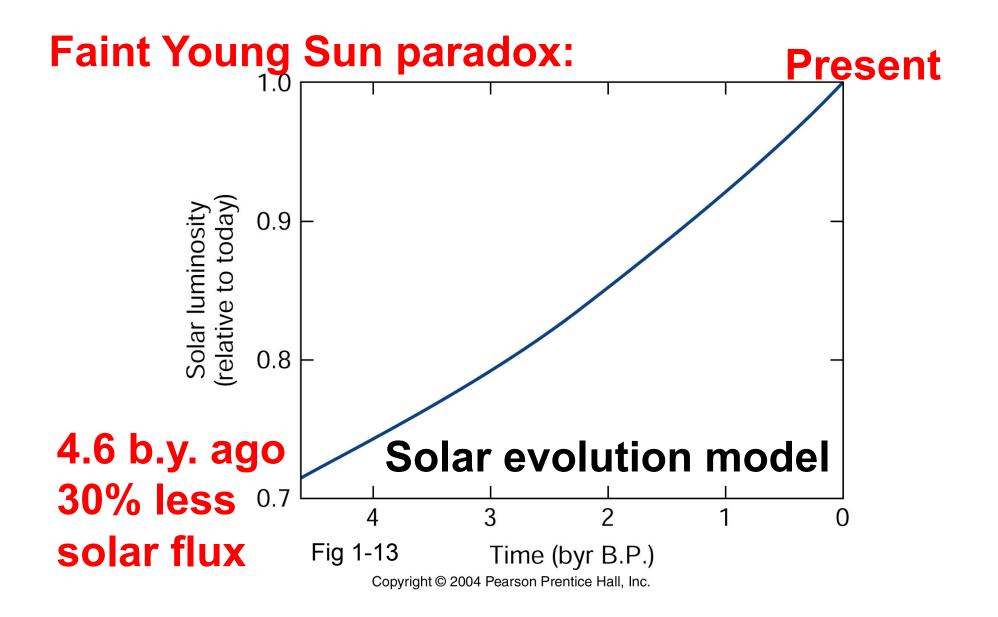
#### ATOC 1060-002 OUR CHANGING ENVIRONMENT

**Class 20 (Chp 12)** 

**Objectives of Today's Class:** 

- 1. The long-term climate regulation;
- 2. The long-term climate record.

### 1. Long-term climate regulation



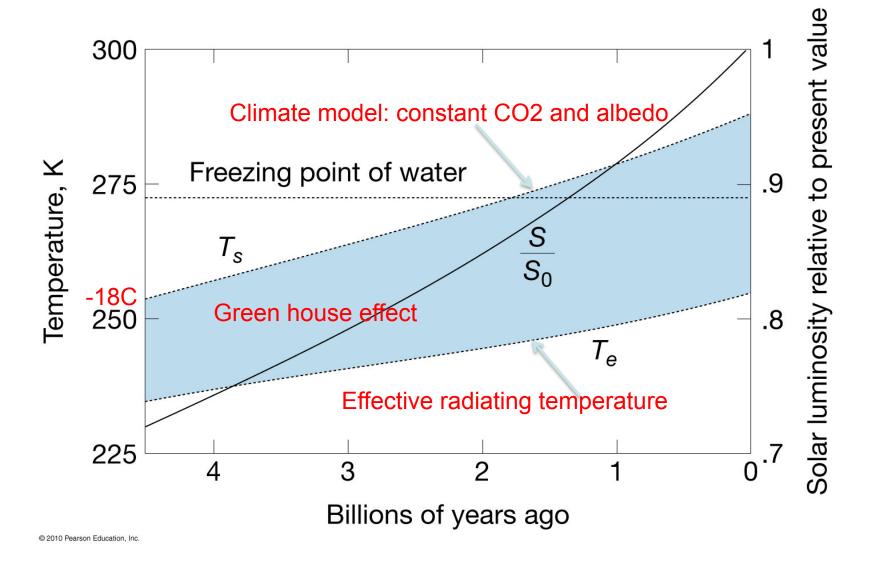


Fig. 12-2

Below freezing: 1.9b.y. ago => contradict geological evidence: liquid water was present 3.8b.y. ago! The earth was warmer than predicted by the fain young sun!

#### Possible solutions:

- i. Lower planetary albedo in the past (0?);
- ii.Additional heat sources besides the Sun (geothermal heat) too small;
- iii.Larger greenhouse effect.

# iii: Larger greenhouse effects (most likely): CO<sub>2</sub>&CH<sub>4</sub>

```
Young Earth, Higher CO<sub>2</sub>: Impact Degassing (carbonate rocks);
Smaller Continents => reduce carbonate rocks storage, reduce silicate rocks weathering=>increase atmospheric CO<sub>2</sub>;
```

```
[Silicate weathering (example): remove atmosphere CO<sub>2</sub>:

CaSiO<sub>3</sub> + 2H<sub>2</sub>CO<sub>3</sub> => Ca<sup>2+</sup> + 2HCO<sub>3</sub> - +SiO<sub>2</sub> + H<sub>2</sub>O]

Wollastonite bicarbonate ion silica

CO<sub>2</sub> dissolve in raindrops (acid)

Silicate weathering process: proportional to temperature;
```

Models: Methane 1000ppm (600times of today, 1.6ppm), 3.8-2.3b.y. ago: Archean Eon (when  $O_2$  levels low)

Favor production of CH<sub>4</sub> (by methanogens)

### A pink sky during Archean?

The Earth: blue sky The Mars larger particles





Scattering blue lights: O<sub>2</sub>,N<sub>2</sub> Archean: CH<sub>4</sub> & CO<sub>2</sub>, polymerize=> bigger long-chains=>scattering orange,red light;

### A pink sky during Archean?

The Earth: blue sky

The Mars: pink (red);

Scattering blue lights: O<sub>2</sub>,N<sub>2</sub>

Archean: CH<sub>4</sub> & CO<sub>2</sub>, polymerize=> bigger long-chains (haze) => scattering orange,red light;

### Climate regulation by the antigreenhouse effect

Was the Earth getting hotter and hotter until no life could survive? NO.

#### **Anti-greenhouse effect:**

CH<sub>4</sub> and haze: strong absorbers of visible (red) light and near-infrared => reradiating back into space without reaching the Earth's surface => cools Earth's surface.

## If haze layer too thick => The Earth too cold

- => CH4 producing bacterial died off
- => reduce CH4
- => thinner haze layer
- => increase Temprature.

Regulating the climate in Archean Era.

### 2. The long-term climate record

Up to now: focus on very early Earth & processes may have contributed to climate stabilization.

Geological indicators=>paleoclimate (past climate) complex => long-term warmth periods & short, intense cold periods, there may have been "Snowball Earth" episodes => suggesting other factors may affect climate as well.

### Paleoclimate:geological indicators

Recent Earth history (millions yrs): estimate ocean temperatures by oxygen isotopes in carbonate sediments from deep-sea cores. [1]C<sub>a</sub>CO<sub>3</sub> => <sup>16</sup>O and <sup>18</sup>O; the colder the water, the more <sup>18</sup>O to be incorporated by minerals ⇒Glacial-interglacial cycles in about 200m.y;

[2] 540m.y. fossil record; species of plants and animals live in certain climates => estimate local surface temperature

### Evidence of past glaciation

Billion-year timescale: Geologic

deposit formed by glacial ice.

**Debris when** glaciers grind up surface rocks=>carried by

glaciers & deposited in piles of rubble -moraines (icesheet margin) 2004 Pearson Prentice Hall, Inc.

**Tillites** 

Fig 12-7a:

### Rocks with long, parallel scratches: glacial striations: moving glaciers drag other rocks across their surface

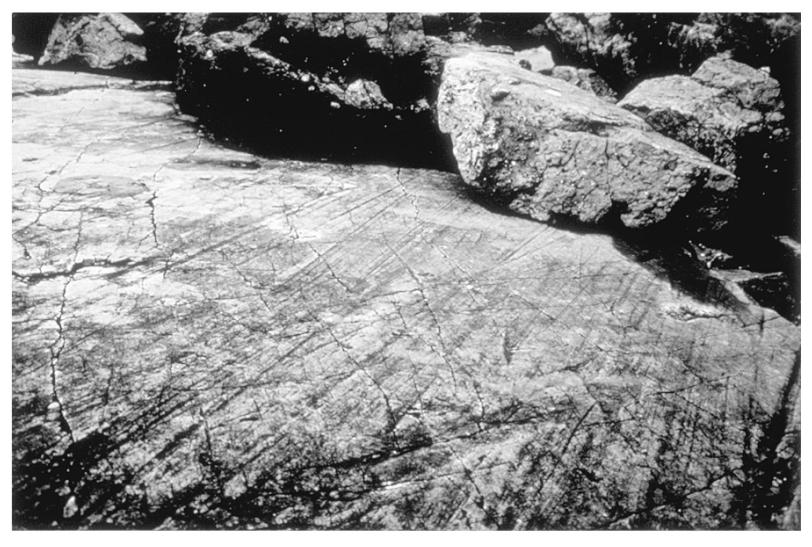


Fig 12-7b (b)
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# "Misplaced" trunks of rock in otherwise finely laminated marine sediments

Rocks
Trapped
In glacial
Ice carried
to sea by
icebergs
Dropstone

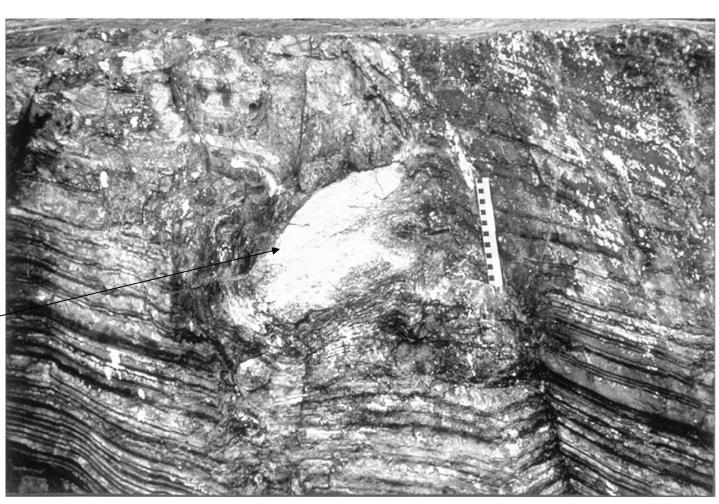


Fig 12-7c (c)
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The long-term glacial record

**Geologists:** 

Earth's

2) 600-800m.y.

climate

Late

history:

**Proterozoic** 

5 main

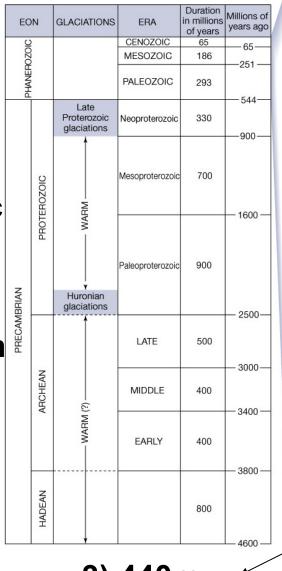
glaciation 1)2.3b.y.

periods.

Huronian

4) ~286m.y. **Permo-Carboniferous** 

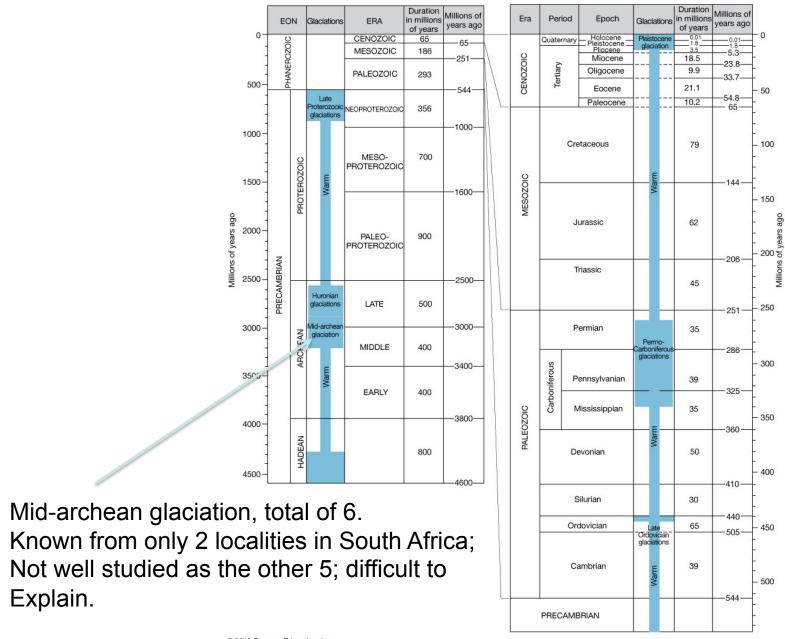
5) ~1.8-recent, **Pleistocene** 



3) 440m.y. Late Ordovician

Duration Millions of Era Period Epoch Glaciations in millions years ago of years Pleistocene Holocene Quaternary Pliocene glaciations 5.3 23.8 Eocene 10.2 Paleocene 65 Cretaceous MESOZOIC 144-Jurassic 62 206-Triassic - 251-Permian Permo-Carboniferous 286-Pennsylvanian 325-Mississippian 360 PALEOZOIC Devonian 410 -Silurian 30 440-Ordovician Ordovician 65 alaciations - 505 -Cambrian 39 544

Textbook 2<sup>nd</sup> edition



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2.3b.y. ago: derived by tillite & dropstone; First found in lake Huron (N. America) => Huronian glaciation. Followed by 1b.y. ice-free conditions.

Why was there a glaciation?

Suppose  $CH_4$  was high in Late Archean (3-2.5b.y) =>  $O_2$  rise around 2.3b.y. due to biological activities (photosynthesis) => eliminate  $CH_4$  => cold! (Geological evidences agree with this).

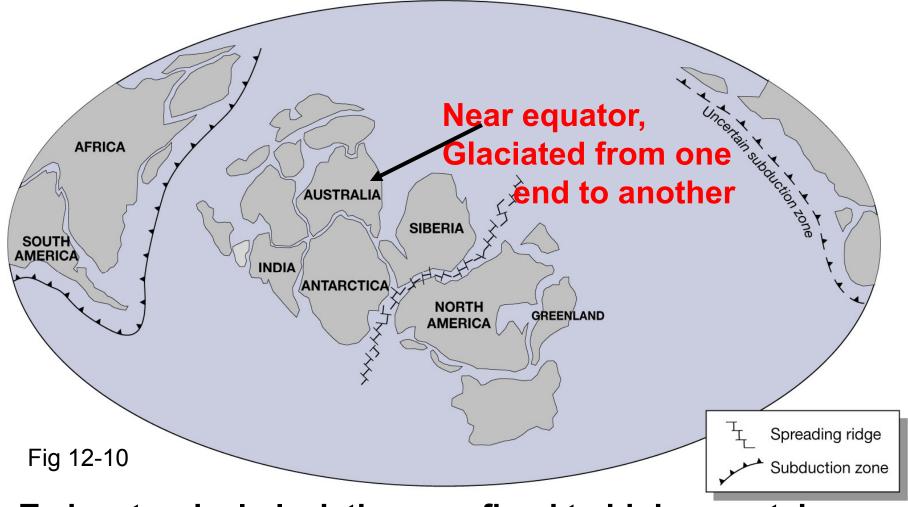
How did the Earth's temperature rise again? Silicate weathering decreases with decrease of temperature => increase CO<sub>2</sub> => increase temperature.

# Low-latitude glaciation: the snowball Earth

The climate became cool once again. Glaciation in the Late Proterozoic (800-600 m.y. ago). [Rising  $CO_2$  => increase Temperature=> increase silicate weathering => decrease  $CO_2$  => cool the temperature.]

Geological evidence: tillites, glacial striations, dropstone were found on 6 of the 7 present day continents (except Antarctica, largely buried by ice); => Snowball Earth.

#### The continents reconstruction at Late proterozoic



Today: tropical glaciation: confined to high mountains; Example: Andes mountains in S. America, above 5km. Geologists: convinced Late proterozoic glaciation is real - all data;